



# Classification of summertime synoptic patterns in Beijing and their associations with boundary layer structure affecting aerosol pollution

Yucong Miao<sup>1</sup>, Jianping Guo<sup>1</sup>, Shuhua Liu<sup>2</sup>, Huan Liu<sup>1</sup>, Zhanqing Li<sup>3,4</sup>, Wanchun Zhang<sup>1</sup>, and Panmao Zhai<sup>1</sup>

<sup>1</sup>State Key Laboratory of Severe Weather & Key Laboratory of Atmospheric Chemistry of CMA, Chinese Academy of Meteorological Sciences, Beijing 100081, China

<sup>2</sup>Department of Atmospheric and Oceanic Sciences, Peking University, Beijing 100871, China

<sup>3</sup>State Key Laboratory of Earth Surface Processes and Resource Ecology & College of Global Change and Earth System Science, Beijing Normal University, Beijing 100875, China

<sup>4</sup>Dept. of Atmospheric & Oceanic Sciences and ESSIC, University of Maryland, College Park, MD 20740, USA

Correspondence to: Jianping Guo (jpuocams@gmail.com) and Zhanqing Li (zli@atmos.umd.edu)

Received: 24 November 2016 – Discussion started: 28 November 2016

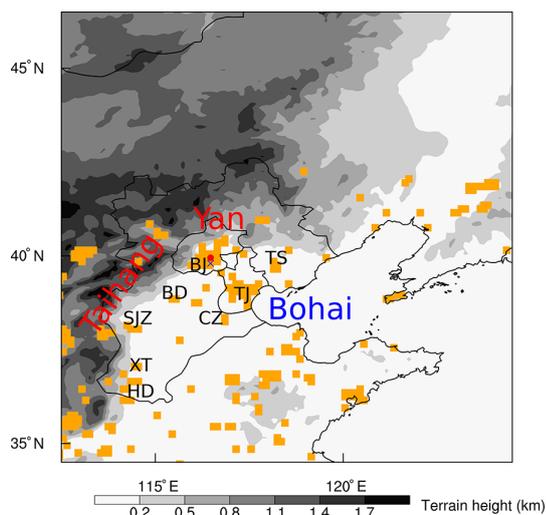
Revised: 1 February 2017 – Accepted: 13 February 2017 – Published: 28 February 2017

**Abstract.** Meteorological conditions within the planetary boundary layer (PBL) are closely governed by large-scale synoptic patterns and play important roles in air quality by directly and indirectly affecting the emission, transport, formation, and deposition of air pollutants. Partly due to the lack of long-term fine-resolution observations of the PBL, the relationships between synoptic patterns, PBL structure, and aerosol pollution in Beijing have not been well understood. This study applied the obliquely rotated principal component analysis in T-mode to classify the summertime synoptic conditions over Beijing using the National Centers for Environmental Prediction reanalysis from 2011 to 2014, and investigated their relationships with PBL structure and aerosol pollution by combining numerical simulations, measurements of surface meteorological variables, fine-resolution soundings, the concentration of particles with diameters less than or equal to 2.5  $\mu\text{m}$ , total cloud cover (CLD), and reanalysis data. Among the seven identified synoptic patterns, three types accounted for 67 % of the total number of cases studied and were associated with heavy aerosol pollution events. These particular synoptic patterns were characterized by high-pressure systems located to the east or southeast of Beijing at the 925 hPa level, which blocked the air flow seaward, and southerly PBL winds that brought in polluted air from the southern industrial zone. The horizontal transport of pollutants induced by the synoptic forcings may be the most im-

portant factor affecting the air quality of Beijing in summer. In the vertical dimension, these three synoptic patterns featured a relatively low boundary layer height (BLH) in the afternoon, accompanied by high CLD and southerly cold advection from the seas within the PBL. The high CLD reduced the solar radiation reaching the surface, and suppressed the thermal turbulence, leading to lower BLH. Besides, the numerical sensitive experiments show that cold advection induced by the large-scale synoptic forcing may have cooled the PBL, leading to an increase in near-surface stability and a decrease in the BLH in the afternoon. Moreover, when warm advection appeared simultaneously above the top level of the PBL, the thermal inversion layer capping the PBL may have been strengthened, resulting in the further suppression of PBL and thus the deterioration of aerosol pollution levels. This study has important implications for understanding the crucial roles that meteorological factors (at both synoptic and local scales) play in modulating and forecasting aerosol pollution in Beijing and its surrounding area.

## 1 Introduction

Beijing, located on the North China Plain (Fig. 1), is the center of politics, culture, and economics in China. With its rapid urbanization, tremendous economic development, and



**Figure 1.** Spatial distribution of the terrain height of the North China Plain. Superimposed are the locations of the US Embassy air quality station ( $39.95^{\circ}$  N,  $116.47^{\circ}$  E, the red dot) and the Beijing meteorological station ( $39.80^{\circ}$  N,  $116.47^{\circ}$  E, the blue cross). Urbanized areas based on MODIS 2012 data are shown in orange. The locations of Beijing (BJ) and adjacent industrial cities are written in black text and include Tianjin (TJ), Tangshan (TS), Baoding (BD), Cangzhou (CZ), Shijiazhuang (SJZ), Xingtai (XT), and Handan (HD).

concomitant increase in energy usage, heavy air pollution episodes largely caused by high aerosol loading have been frequently reported in Beijing (Chan and Yao, 2008; Guo et al., 2011, 2013; San Martini et al., 2015; Zhang and Cao, 2015). Great efforts, therefore, have been devoted to investigating air quality issues in Beijing through comprehensive observational and modeling studies (He et al., 2001; Zhu et al., 2011; Liu et al., 2013; Quan et al., 2013; Zhang et al., 2013; Guo et al., 2014; Wang et al., 2015a; Z. Zhang et al., 2015; Ye et al., 2016). The major sources of aerosol in Beijing include traffic emission, power plants, industry, domestic emissions, and agricultural activities (R. Zhang et al., 2015; Liu et al., 2016; Peng et al., 2016). In an annual cycle, the aerosol pollution in Beijing has different characteristics in different seasons. In spring, the majority of heavy aerosol pollution is associated with dust storms (He et al., 2001; Guo et al., 2013). During summer and fall, photochemical production and agricultural burning may play a role in exacerbating the air quality (Zhang et al., 2013; Wang et al., 2015b). In winter, frequent severe haze events were found to be associated with the substantial increase in coal combustion for heating (Zhang et al., 2013; Tang et al., 2016). Air quality can be further deteriorated through secondary aerosol formation (e.g., Huang et al., 2014; Han et al., 2015).

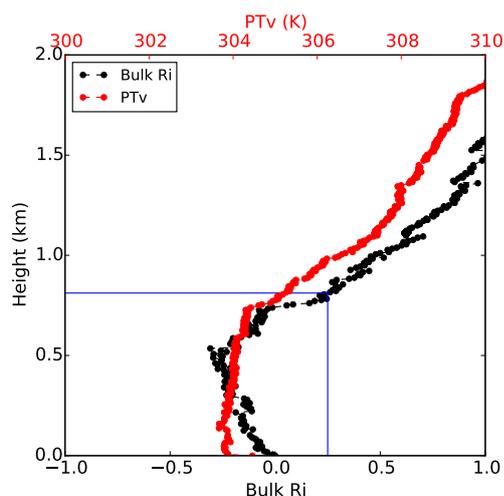
In addition to high emissions and aerosol chemistry processes, meteorological conditions also play important roles in the formation and evolution of aerosol pollution in Beijing (Zhang et al., 2012; Quan et al., 2013; Hu et al., 2014;

Miao et al., 2015b; Ye et al., 2016; Z. Zhang et al., 2015). In the heavily populated monsoon region (e.g., Beijing), the large-scale distributions and variations of aerosol loading are strongly influenced by the monsoon circulations (Zhang et al., 2010; Li et al., 2016; Wu et al., 2016), since the monsoon circulations directly determine the large-scale transport and lifetime of aerosols. In addition to the increase in anthropogenic emissions during the past decades in China, the recently frequent occurrence of aerosol pollution in China has been found to be associated with the variability in monsoon (Niu et al., 2010; Zhang et al., 2010; Liu et al., 2011).

On a regional scale, detailed statistical analyses have also shown that heavy aerosol pollution in Beijing is associated with southerly winds, high relative humidity (RH), stable atmospheric stratification, and a low boundary layer height (BLH) (Quan et al., 2013; H. Zhang et al., 2015; Tang et al., 2016; Ye et al., 2016). Among these local meteorological factors, the BLH is one of the most crucial factors because it directly determines the total dispersion volume (Quan et al., 2013; Hu et al., 2014; Tang et al., 2016) and has a strong influence on the occurrence, maintenance, and dissipation of aerosol pollution in Beijing (e.g., Miao et al., 2015b).

With the Yan and Taihang mountains to the north and west of Beijing (Fig. 1), the thermally induced mountain–plain breeze circulation can be well developed under favorable synoptic conditions, which impacts on the PBL structure and modulates aerosol pollution (De Wekker, 2008; Chen et al., 2009; Liu et al., 2009; Hu et al., 2014; Miao et al., 2015b, 2016). Since the Bohai Sea is located  $\sim 150$  km to the southeast of China (Fig. 1), a sea breeze (Miller, 2003) can be established and penetrate inland to Beijing. This can also play a role in affecting the PBL structure and air quality (Liu et al., 2009; Sun et al., 2013; Miao et al., 2015a, b). The diurnal variation of land breeze and sea breeze provides a mechanism for the pollutants in the Beijing–Tianjin–Hebei region to be recirculated and accumulated; in the evening and early morning, the presence of land breeze (offshore wind) could bring the pollutants emitted from coastal regions to the Bohai Sea and, then, in the afternoon, the development of sea breeze (onshore wind) could bring these pollutants back to coastal regions, leading to exacerbation of pollution. When the sea breeze penetrates further inland, the pollutants emitted from coastal regions could be transported to the downstream regions.

Although the importance of the PBL structure for air pollution has been widely recognized (Wang et al., 2014; Miao et al., 2015b; Tang et al., 2016), more investigations are warranted concerning (1) the crucial factors affecting the development of the PBL and (2) the relationships between large-scale synoptic forcing and the structure of the PBL. They are yet to be fully understood, partly due to the lack of long-term fine-resolution PBL observations (Liu and Liang, 2010). Radiosondes are conventionally launched twice a day at 00:00 (08:00) and 12:00 (20:00) Coordinated Universal Time (UTC) (Beijing Local Time, BJT). Most data are only



**Figure 2.** Vertical profiles of the bulk Richardson number ( $Ri$ , in black) and virtual potential temperature (PTv, in red) based on L-band sounding observations made in Beijing on 30 June 2013 at 14:00 BJT. The boundary layer height is the height where  $Ri$  first reaches the value of 0.25 (the blue lines).

reported at significant pressure levels with at most six records below 500 hPa (Liu and Liang, 2010), which cannot capture the fine structure of the PBL.

In 2011, an L-band radiosonde network across China (Guo et al., 2016; Zhang et al., 2016) was established by the China Meteorological Administration. This network of measurements provides fine-resolution profiles of temperature (Fig. 2), pressure, RH, and wind speed and direction twice a day (08:00 and 20:00 BJT). Additional soundings are made at 14:00 BJT in the summertime (June–July–August, the wet season) at the Beijing site (39.80° N, 116.47° E) (the blue cross in Fig. 1). These fine-resolution sounding observations allow for the investigation of the PBL structure over Beijing.

The development of the PBL is mainly controlled by large-scale external synoptic conditions and by local surface sensible heat fluxes (Garratt, 1994). So, in this study, not only are local factors that affect the surface heat budget analyzed (e.g., cloud cover), but also the large-scale synoptic forcing. A synoptic regime is defined by large-scale warm/cold advection and transport pathways of water vapor and pollutants (Zhang et al., 2012; Zhao et al., 2013; Ye et al., 2016), which can affect the PBL structure and air quality (Miao et al., 2015b; Ye et al., 2016). The variation of the synoptic patterns modulated the ambient pollutants and likely provided the primary driving force for the day-to-day variations in air pollution level (e.g., Chen et al., 2008; Wei et al., 2011; Zhang et al., 2012; Hu et al., 2014; Ding et al., 2016). Besides, severe aerosol pollution events in Beijing tend to occur more frequently under stable and weak anticyclone synoptic conditions (e.g., Zhang et al., 2012; Liu et al., 2013; Li et al., 2015). However, most previous studies evaluated the impacts of synoptic types through case studies for short periods (Chen et al., 2009; Hu

et al., 2014; Quan et al., 2014; Tie et al., 2015; Miao et al., 2016), which is not suitable for identifying dominant synoptic patterns.

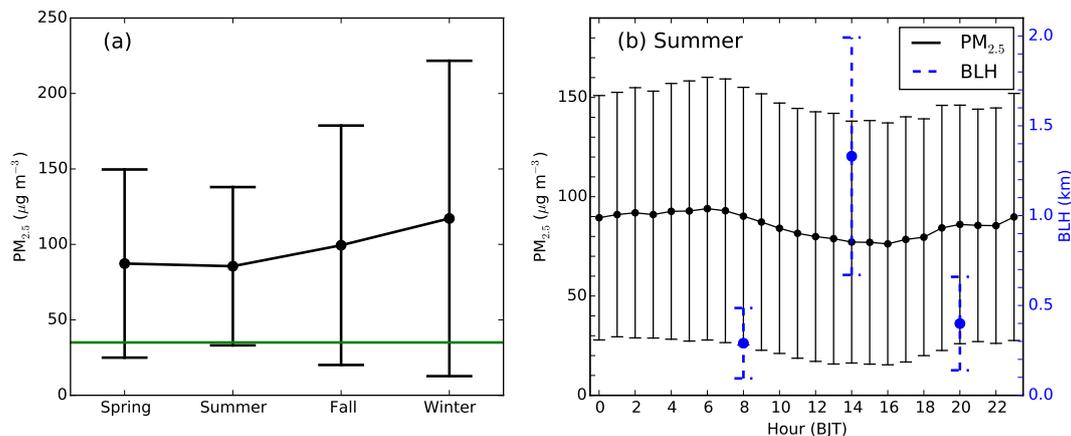
In this study, we employ a climatological approach to classify the summertime synoptic patterns over Beijing from 2011 to 2014, the same objective classification approach applied by Zhang et al. (2012) and Ye et al. (2016) to classify synoptic types in the North China Plain. Zhang et al. (2012) identified nine synoptic patterns using the daily surface-level pressure fields from 2000 to 2009, and Ye et al. (2016) presented a more detailed classification of synoptic patterns for fall and winter. They found that the poor air quality in Beijing was associated with high pressure to the east and a relatively low BLH. However, previous studies did not unravel the causes of the low BLH and the physical mechanisms underlying it, partly due to the lack of appropriate observational data (e.g., fine-resolution soundings in the afternoon, cloud cover). In addition, the north–south movement of the subtropical anticyclone plays an important role in modulating the seasonal variation in prevailing synoptic patterns in Beijing (Miao et al., 2015b). It is thus better to classify synoptic patterns for each season.

In summer, although the aerosol pollution level in Beijing is lower than that in fall and winter (Fig. 3a), the seasonally averaged concentration of particles with diameters less than or equal to 2.5  $\mu\text{m}$  ( $\text{PM}_{2.5}$ ) is still as high as 85.7  $\mu\text{g m}^{-3}$ , based on data from 2011 to 2014. This is 2.4 times higher than the national standard level (35  $\mu\text{g m}^{-3}$ ). Also, the relationships between aerosol pollution, the PBL structure, and synoptic patterns in summer are rarely studied, which will be systematically evaluated using long-term observations. This study will extend previous studies as it is an attempt to understand the impacts of large-scale synoptic forcing on the PBL structure and air quality. The remainder of this paper is organized as follows. In Sect. 2, the methodology and data are described. In Sect. 3, the summertime synoptic types are classified, and their relationships with aerosol pollution and the PBL structure are investigated. The main findings are summarized in Sect. 4.

## 2 Data and methods

### 2.1 Data

To classify the summertime synoptic types in Beijing, geopotential height (GH) fields derived from the National Centre for Environmental Prediction (NCEP) global Final (FNL) reanalysis (<http://rda.ucar.edu/datasets/ds083.2/>) from 2011 to 2014 are used. The NCEP-FNL reanalysis is produced by the Global Data Assimilation System, which continuously assimilates observations from the Global Telecommunication System and other sources. The NCEP-FNL reanalysis fields are on  $1^\circ \times 1^\circ$  grids with a 6 h temporal resolution,



**Figure 3.** (a) Seasonal variation of the daily averaged PM<sub>2.5</sub> concentration (the mean  $\pm$  1 standard deviation), and (b) diurnal cycle of the PM<sub>2.5</sub> concentration (the mean  $\pm$  1 standard deviation) in summer derived from hourly measurements made at the US Embassy air quality station from 2011 to 2014. The green line in (a) shows the national standard level ( $35 \mu\text{g m}^{-3}$ ). The diurnal cycle of the BLH (in blue, the mean  $\pm$  1 standard deviation) derived from summertime soundings in Beijing is shown in (b) on the right-hand ordinate.

i.e., 00:00 (08:00), 06:00 (14:00), 12:00 (20:00), and 20:00 (02:00) UTC (BJT).

In this study, daily GH fields at the 925 hPa level from the NCEP–FNL reanalysis covering the Beijing–Tianjin–Hebei (BTH) region ( $43\text{--}47^\circ\text{N}$ ,  $112\text{--}125^\circ\text{E}$ ) (Fig. 1) were classified to identify the prevailing synoptic types in summer. Results from the classification of the 925 hPa GH field are similar to those using the GH fields at other tropospheric levels because there is a high degree of dependence among individual levels (Huth et al., 2008).

To investigate the PBL structures associated with different synoptic types, summertime soundings collected at the Beijing site ( $39.80^\circ\text{N}$ ,  $116.47^\circ\text{E}$ ) for the period 2011–2014 were analyzed. This sounding station, equipped with an L-band radiosonde system (Guo et al., 2016), provides atmospheric sounding data (profiles of temperature, RH, wind speed and direction) up to three times a day (08:00, 14:00, and 20:00 BJT) at a high vertical resolution. In total, 1055 effective soundings were obtained for this study: 357 soundings at 08:00 BJT, 332 soundings at 14:00 BJT, and 366 soundings at 20:00 BJT. In addition, the total cloud cover (CLD) was observed four times a day (02:00, 08:00, 14:00, and 20:00 BJT) and hourly near-surface observations (temperature, RH, wind speed and direction, and precipitation amount) were made.

The aerosol pollution level in Beijing is denoted by the near-surface PM<sub>2.5</sub> concentration. Since 2008, US diplomatic missions in China have monitored PM<sub>2.5</sub> concentrations and have made both real-time and historic data available to the public ([www.stateair.net](http://www.stateair.net)). Hourly PM<sub>2.5</sub> concentrations in Beijing (the red dot in Fig. 1) are measured using a beta attenuation monitor (BAM) (Chung et al., 2001) installed on the roof of the US Embassy ( $39.95^\circ\text{N}$ ,  $116.47^\circ\text{E}$ ). The BAM technique is a reference method for measuring PM<sub>2.5</sub> concentrations that is used by the US Environmental

Protection Agency. In this study, summertime PM<sub>2.5</sub> concentration measurements from 2011 to 2014 were used to investigate the relationship between aerosol pollution and synoptic patterns in Beijing.

## 2.2 BLH derived from soundings

The bulk Richardson number ( $Ri$ ) method (Vogelezang and Holtslag, 1996) was applied to estimate the BLH in Beijing from sounding data because it is suitable for both stable and convective PBLs (Seidel et al., 2012). The  $Ri$  is defined as the ratio of turbulence associated with buoyancy to the turbulence caused by mechanical shear:

$$Ri(z) = \frac{(g/\theta_{vs})(\theta_{vz} - \theta_{vs})(z - z_s)}{(u_z - u_s)^2 + (v_z - v_s)^2 + bu_*^2}, \quad (1)$$

where  $z$  is the height (above ground level, a.g.l.),  $g$  is the acceleration caused by gravity,  $\theta_v$  is the virtual potential temperature,  $u$  and  $v$  are the components of the observed wind speed,  $b$  is a constant, and  $u_*$  is the surface friction velocity. The subscript “s” denotes the surface level. Since  $u_*$  is not known from the sounding observations, and its magnitude is much smaller than that of the bulk wind shear term in the denominator (Vogelezang and Holtslag, 1996), we set  $b = 0$  and ignore the surface frictional effect. The BLH is referred to as the lowest level  $z$  at which the interpolated  $Ri$  crosses the critical value of 0.25. A similar criterion was applied to investigate PBL climatologies by Seidel et al. (2012) for the US and by Guo et al. (2016) for China. A case in point for the BLH derived from the sounding profiles of  $\theta_v$  and  $Ri$  at the Beijing site is shown in Fig. 2.

## 2.3 Classification of synoptic types

Approaches used to classify synoptic patterns can be roughly split into two groups: subjective and objective. The objec-

tive approach is usually referred to as manual classification, which subjectively defines a priori and in which the case assignment is also subjective (Huth et al., 2008). A subjective classification is arbitrary to a large extent. By contrast, an objective approach defines types and assigns cases using numerical procedures based on the measures of (dis)similarity and variance maximization. Because an objective classification is capable of processing large amounts of data and depends less on one's experience, we choose the objective approach (Huth et al., 2008) to identify synoptic types in Beijing. Studying large-scale synoptic patterns allows us to consider the numerous interrelated meteorological variables within an integrated framework (Zhang et al., 2012), thus providing an insight into the physical mechanisms underlying aerosol pollution in Beijing.

The obliquely rotated principal component analysis in the T-mode (T-PCA) approach (Richman, 1981; Huth et al., 2008) was first used to analyze large-scale synoptic conditions through classification of the predominant synoptic types, which was adopted for air quality studies (e.g., Stefan et al., 2010; Zhang et al., 2012). The T-PCA calculates the eigenvectors of the input data set by singular value decomposition and finds typical patterns by loadings that can be divided into classes. The application of the PCA in T-mode means that daily patterns form the columns in the input data matrix and grid-point values form its rows (Huth, 2000). Unlike the common application of PCA in the S-mode, which is used to isolate subgroups of grid points that co-vary similarly, the T-PCA is used to isolate subgroups of similar spatial patterns. This approach has proven to be a reliable classification method, largely due to its temporal and spatial stability, in addition to its ability to reproduce predefined dominant patterns with little dependence on pre-set parameters (Huth, 1996; Philipp et al., 2010).

In this study, the T-PCA classification based on Huth (2000) was done using the *cost733class* software package (<http://cost733.met.no>), which was developed for creating, comparing, and evaluating classifications in several variants. To speed up the calculation of PCA, the data are split into 10 subsets, and then the principle components (PCs) obtained from each subset are projected onto the remaining data. The T-PCA classification based on the *cost733* software package includes the following steps.

1. The data are standardized spatially. Each pattern's mean is subtracted from the data, and then the patterns are divided by their standard deviations.
2. The data are split into 10 subsets by selecting the data once every 10 days. For example, the first subset consists of the 1st, 11th, 21st, 31st, etc. days, and the second subset consists of the 2nd, 12th, 22nd, 32nd, etc. days.
3. The PCs are calculated using the singular value decomposition for each subset. And the PCs of each subset are

ordered according to the magnitude of their explained variances.

4. An oblique rotation (using direct oblimin) is applied to the PCs, employing an adaptation of the Gradient Projection Algorithm of Bernaards and Jennrich (2005). The main reason for using rotation is to facilitate the interpretation (Abdi and Williams, 2010). This transformation does not constrain the orthogonality, allowing for the PCs the freedom to better reflect the original data (Richman, 1981).
5. The PC scores of each subset are projected onto the remaining data by solving the matrix equation  $\Phi A^T = F^T Z$ , where  $F$  and  $\Phi$  are matrices of PC scores and PC correlations, respectively,  $Z$  is the full data matrix, and  $A$  are pseudo-loadings to be determined. Each day is classified with the PC (type) for which it has the highest loading.
6. Contingency tables are finally used to compare the 10 classifications, and the classification most consistent with the other 9 classifications is selected as the resultant one.

Using the T-PCA module of the *cost733class* software package to classify synoptic types, the number of PCs needs to be explicitly defined. To understand the relationships between synoptic types, PBL structures, and aerosol pollution, we tested the classification using different numbers of principal components (e.g., 4, 5, 6, 7, 8, 9), and compared the results with the observed BLH and PM<sub>2.5</sub> concentration (Fig. S1 in the Supplement). The classification using seven principal components shows the most significant anti-correlation between BLH and PM<sub>2.5</sub> concentration ( $R = -0.97$ ,  $p < 0.01$ , Fig. S1). This will be discussed further in Sect. 3.

## 2.4 Idealized numerical experiments

Considering the PBL structure being simultaneously modulated by the cloudiness, warm/cold advection, and aerosols, to better understand the relationships between the synoptic condition and PBL structure in Beijing, the Weather Research and Forecasting model (WRF version 3.7.1) was used to conduct seven idealized numerical experiments: each corresponds to one of the seven identified synoptic patterns.

In the idealized experiments, the simulation region was set as the same studied region shown in Fig. 1, with a horizontal grid spacing of 0.1° (~11 km). In the vertical dimension, 48 layers were set from the surface to the 100 hPa level, with 21 layers between the surface and 2 km above ground level (a.g.l.). To isolate the effects of synoptic forcing (e.g., warm/cold advection) from cloudiness, the microphysics and cumulus schemes were turned off in the idealized simulations. Similar approaches have been used by De Wekker (2008) to investigate the suppression of BLH near a

mountain, and by Pu and Dickinson (2014) to investigate the dynamics of a low-level jet over the Great Plains. The other physics parameterization schemes utilized by this study included the Yonsei University (YSU) PBL scheme (Hong et al., 2006), the updated rapid radiative radiation scheme (Iacono et al., 2008), and the Noah land surface scheme (Chen and Dudhia, 2001) with a single-layer urban canopy model (Kusaka et al., 2001).

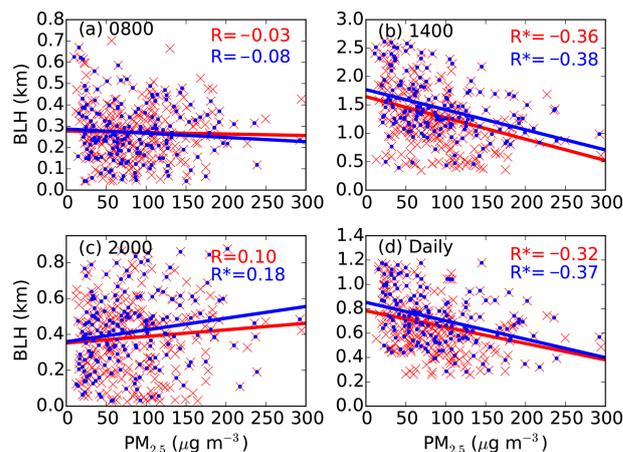
For the initial and boundary conditions (IBCs) of humidity and temperature for the atmosphere, ocean, and soil, all seven idealized simulations were set up using the summertime seasonal means derived from the NCEP-FNL reanalysis from 2011 to 2014. And the IBCs of wind and pressure were set up using the means of each synoptic pattern obtained from the NCEP-FNL reanalysis. All these boundary conditions were cycled periodically in time, i.e., from 00:00 to 06:00 to 12:00 to 18:00 UTC, and then back to 00:00 UTC. To represent the seasonal mean solar radiation forcing in summer, all seven of these simulations were initialized at 08:00 BJT (00:00 UTC), 14 July. The simulations were run for 42 h, and the first 30 h were considered as the spin-up period. Such numerical configurations allowed us to examine the influences of different synoptic forcings on the PBL structure in Beijing, isolating the effects of cloudiness.

### 3 Results and discussion

#### 3.1 Relationships between the BLH and aerosol pollution in summer

During a diurnal cycle, the development of the PBL is closely tied to solar heating of the ground (Stull, 1988). After sunrise, the PBL in Beijing during summer undergoes a transition from a nocturnal stable PBL to a convective PBL, and reaches its maximum depth in the afternoon (cf. Fig. 3b), remaining there until sunset (Stull, 1988). After sunset, without sufficient heat fluxes to maintain the convective PBL, the BLH drops quickly (Fig. 3b). Along with the diurnal evolution of the BLH, the near-surface  $\text{PM}_{2.5}$  concentration exhibits a nearly reversed diurnal variation. The  $\text{PM}_{2.5}$  concentration is generally low in the afternoon and reaches its peak values in the evening and early morning.

As shown in Fig. 4, the BLHs derived from soundings were compared with their corresponding daily average  $\text{PM}_{2.5}$  concentrations. At 08:00 BJT, the PBL is typically shallow and its depth is statistically uncorrelated with the daily  $\text{PM}_{2.5}$  concentration ( $R = -0.03$ ,  $p = 0.62$ , Fig. 4a). At 14:00 BJT, the PBL is fully developed, with an average BLH of  $\sim 1.3$  km, which is clearly anti-correlated with the daily  $\text{PM}_{2.5}$  concentration ( $R = -0.36$ ,  $p < 0.01$ , Fig. 4b). This implies that the variation in afternoon BLH plays an important role in modulating the day-to-day variation of aerosol pollution level in Beijing. At 20:00 BJT, the PBL is in transition from a convective state to a nocturnal stable state, and



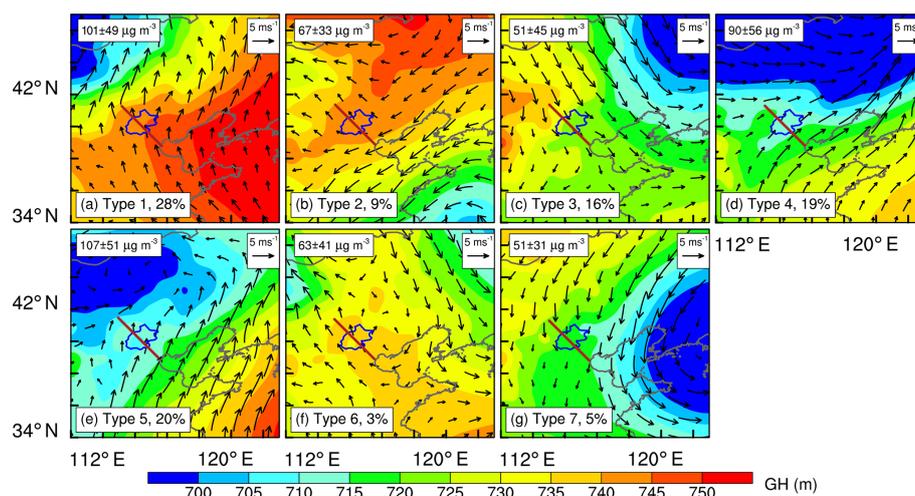
**Figure 4.** The BLH as a function of daily  $\text{PM}_{2.5}$  concentration at (a) 08:00 BJT, (b) 14:00 BJT, (c) 20:00 BJT, and (d) averaged over the whole day (in red). BLHs greater than the 95th percentile value and less than the 5th percentile value are not included in the plots. The correlations between the  $\text{PM}_{2.5}$  concentration and BLH without considering measurements made on rainy days are shown in blue. The correlation coefficients ( $R$ ) are given in each panel, and the asterisks indicate values that are statistically significant ( $p < 0.05$ ). The  $p$  values were calculated based on the Student's  $t$  distribution.

the BLH becomes uncorrelated with the daily  $\text{PM}_{2.5}$  concentration ( $R = 0.10$ ,  $p = 0.09$ , Fig. 4c). Averaged over the whole day, a significantly negative correlation ( $R = -0.32$ ,  $p < 0.01$ ) is still found between the daily BLH and  $\text{PM}_{2.5}$  concentration (Fig. 4d). When excluding observations made on rainy days, a stronger correlation ( $R = -0.37$ ,  $p < 0.01$ ) is obtained. So, in this study, we mainly investigate the factors that determine the BLH at 14:00 BJT and their relationships with large-scale synoptic patterns.

#### 3.2 Synoptic patterns and aerosol pollution

Using the T-PCA classification based on the summertime 925 hPa GH fields, seven dominant types of synoptic patterns (Fig. 5 and Table S1) were identified. According to the locations of high- and low-pressure systems with respect to Beijing (Fig. 5), these seven synoptic types can be briefly described as (1) high pressure to the east, (2) high pressure to the north, (3) low pressure to the northeast, (4) low pressure to the north, (5) low pressure to the northwest, (6) weak high pressure over Beijing, and (7) low pressure to the east.

Among these seven identified synoptic patterns, Types 1, 4, and 5 are associated with heavier aerosol pollution (Fig. 5), all with an average  $\text{PM}_{2.5}$  concentration greater than  $90 \mu\text{g m}^{-3}$ . These three types are also the most frequent synoptic patterns in summer, accounting for 67% of the total. At the 925 hPa level, the three synoptic patterns are characterized by a high-pressure system located to the east or southeast of Beijing, which brings southerly winds to Bei-



**Figure 5.** The 925 hPa geopotential height (GH) fields (colored areas) and wind vector fields (arrows) in summer from 2011 to 2014 for the seven synoptic patterns: (a) Type 1, (b) Type 2, (c) Type 3, (d) Type 4, (e) Type 5, (f) Type 6, and (g) Type 7. The occurrence frequency of each synoptic pattern is given in the bottom left of each panel and the  $\text{PM}_{2.5}$  concentration (the mean  $\pm 1$  standard deviation) is shown in the top left of each panel. The location of the Beijing metropolitan area is outlined in blue near the center of each panel. The red lines cutting through the Beijing metropolitan area in (a)–(g) indicate the location of the cross sections shown in Fig. 11.

jing (Figs. 5 and S2). Southerly winds are not only observed at the 925 hPa level, but also throughout almost all of the PBL (Fig. 6a). With the southerly PBL winds blowing over the plains of the BTH region, pollutants emitted from the surrounding southern cities (e.g., Baoding, Shijiazhuang, and Cangzhou) can be transported to Beijing (Wang et al., 2010; Zhang et al., 2012; Miao et al., 2016), leading to a worsening of the air quality in Beijing. The impacts of synoptic forcings (i.e., high pressure to the east or southeast of Beijing and resultant southerly winds) on the transport of pollutants from adjacent regions to Beijing have also been reported by Zhang et al. (2012) and Ye et al. (2016).

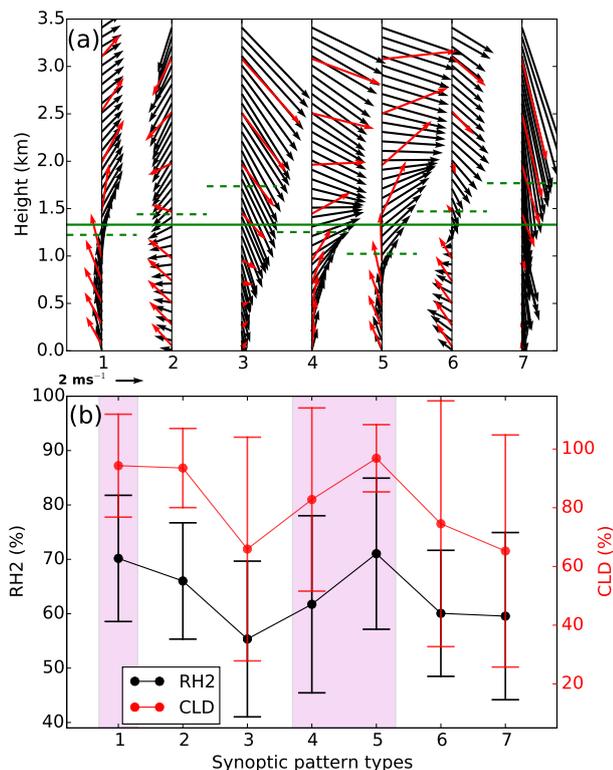
In addition to the southerly PBL winds, high 2 m RH (RH2), high CLD, and low BLH are other crucial parameters associated with severe aerosol pollution (Figs. 6 and 7, Table S2). At 14:00 BJT, all three synoptic patterns (Types 1, 4, and 5) have a relatively low BLH (Fig. 6a and Table S2), with an average value of less than 1.4 km. This would limit the vertical dispersion of pollutants to some extent and exacerbate the pollution. Meanwhile, Types 1 and 5 are characterized by relatively high CLD and RH2 (Fig. 6b), which could facilitate the hygroscopic growth of aerosols (Kim et al., 2014; Ye et al., 2016).

Among all these meteorological factors/processes, the horizontal transport of pollutants driven by different synoptic patterns may be the most important factor modulating the air quality. For example, Fig. 5 reveals a large difference in  $\text{PM}_{2.5}$  concentration between Type 1 ( $101 \mu\text{g m}^{-3}$ ) and Type 2 ( $67 \mu\text{g m}^{-3}$ ), while their CLD and RH2 are quite similar, and the difference in BLH is only  $\sim 0.2$  km (Fig. 6). For these two synoptic types, the large difference in aerosol concentration could be attributed to the different horizontal trans-

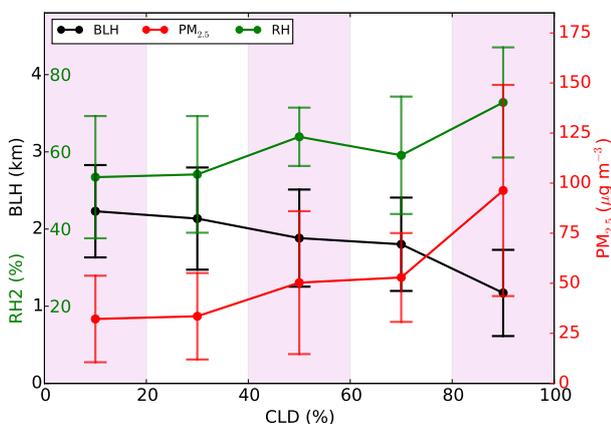
port induced by the different prevailing winds (i.e., southerly wind for Type 1 and easterly wind for Type 2). And the pollution of Type 1 may be exacerbated by the relatively low BLH in the afternoon. Besides, during the pollution episodes, the presence of aerosols within the PBL may also play a role in modulating the PBL structure by portioning the solar radiation between the land surface and the PBL (Li et al., 2007; Ding et al., 2016; Miao et al., 2016). On the one hand, the presence of aerosols could reduce the solar radiation reaching the land surface. The absorbing aerosols could heat the atmosphere, leading to a more stable and shallower PBL during the daytime (Quan et al., 2013; Gao et al., 2015).

Additionally, as illustrated in Fig. 7, the RH2, CLD, and BLH are highly related. The increase in CLD reduces the amount of solar radiation reaching the surface, and suppresses the development of the PBL and the vertical mixing of water vapor, leading to a decrease in the BLH and an increase in RH2. The high CLD could be at least partially responsible for the relative shallow PBL in the afternoon. On the other hand, when RH2 increases, the lifting condensational level (LCL) can drop, favoring the formation of cumulus clouds, which may subsequently suppress the development of the PBL (Wilde et al., 1985; Craven et al., 2002; Zhu and Albrecht, 2002). As a result, under clear conditions ( $\text{CLD} < 20\%$ ), the average RH2 is less than 40%, while the average BLH at 14:00 BJT can reach  $\sim 2.2$  km (Fig. 7), favoring the vertical dispersion of aerosols. By contrast, under cloudy conditions ( $\text{CLD} > 80\%$ ), the average RH2 and BLH increase to  $\sim 70\%$  and decrease to  $\sim 1.2$  km, respectively.

Figure 8 shows the correlations between average  $\text{PM}_{2.5}$  concentrations and meteorological variables for the different synoptic patterns. The  $\text{PM}_{2.5}$  concentration in Beijing is sig-



**Figure 6.** (a) Average wind vector profiles at 14:00 BJT and (b) observed daily RH2 and CLD at 14:00 BJT (mean  $\pm$  1 standard deviation) associated with the seven types of synoptic pattern. In (a), the wind vectors derived from soundings are denoted by black vectors and those from the NCEP-FNL reanalysis are in red. The green solid line shows the seasonal mean BLH and the dashed green lines represent the mean BLH for each synoptic pattern type. Note that the wind vector profiles from the NCEP-FNL reanalysis were derived by interpolating the values from the four nearest grids.



**Figure 7.** The BLH at 14:00 BJT (in black), daily RH2 (in green), and daily  $\text{PM}_{2.5}$  concentration (in red) as a function of CLD at 14:00 BJT. Mean values  $\pm$  1 standard deviation are shown. The bin size is 20 %.

nificantly correlated with the southerly wind at the 925 hPa level, RH2, and CLD at 14:00 BJT, and anti-correlated with the BLH at 14:00 BJT. By contrast, the  $\text{PM}_{2.5}$  concentration is uncorrelated with the 2 m temperature and wind speed. Type 1, 4, and 5 synoptic patterns associated with southerly PBL winds, high CLD, low BLH, and high RH2 in Beijing favor the occurrence of heavy aerosol pollution in summer. By contrast, synoptic patterns with strong northerly PBL winds (Types 3 and 7) or relatively high BLH (Types 2 and 6) are associated with good air quality conditions in Beijing.

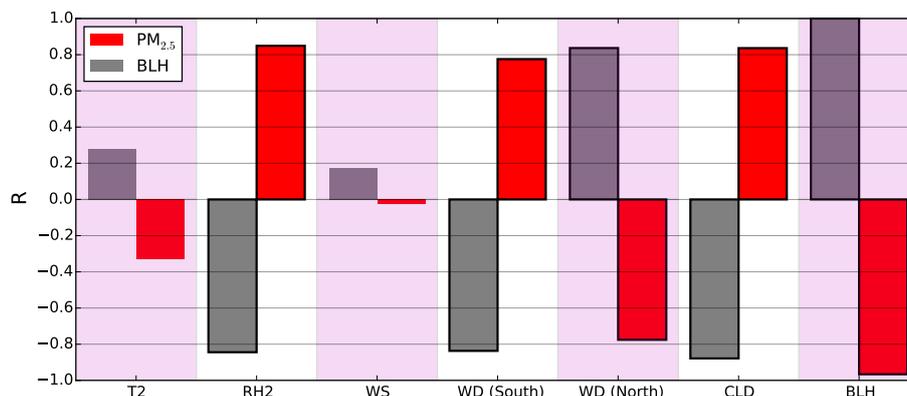
In addition to the aforementioned meteorological factors, precipitation may also affect aerosol pollution levels through wet scavenging (Yoo et al., 2014). Therefore, for each identified synoptic pattern, we examined the aerosol pollution after removing observations made on rainy days (Figs. S3 and S4). Similar relationships between aerosol pollution, RH2, CLD, BLH, and synoptic patterns were found, suggesting that whether it rains or not is not important when it comes to understanding effects of synoptic patterns on PBL structure and aerosol pollution in Beijing.

### 3.3 Large-scale synoptic warm/cold advection and PBL structures

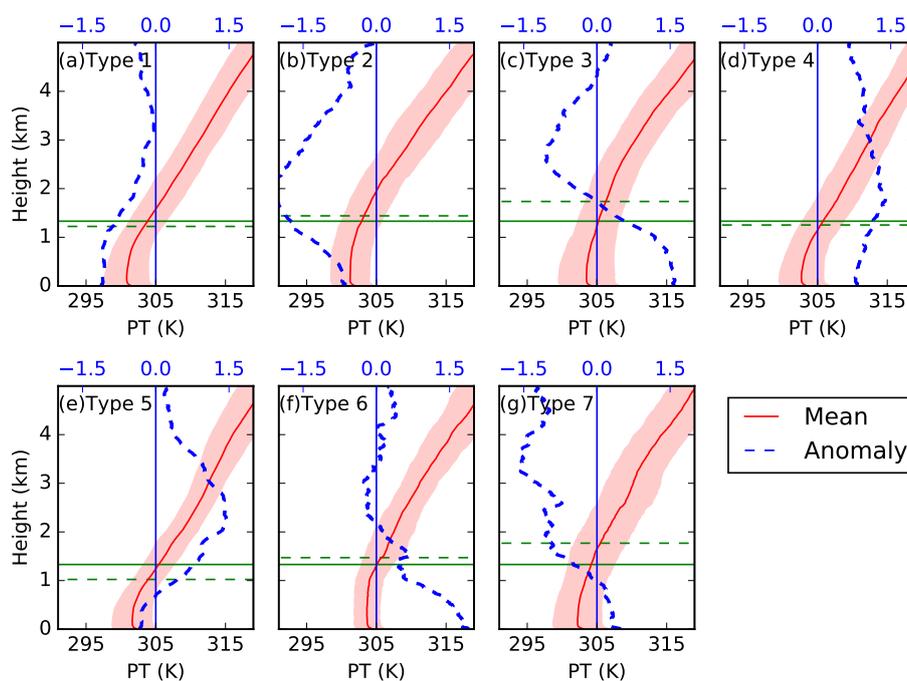
Figure 9 shows the averaged PT profile and its anomaly (subtracted from the seasonally averaged PT profile in summer) for each synoptic type. For Types 1, 4, and 5, the differential PT anomalies within and above the PBL play a role in enhancing the thermal inversion at the PBL top, which would suppress the development of the PBL (Fig. 9a, d, and e). Specifically, the cool (warm) anomaly of Type 1 (4) is stronger (weaker) within the PBL than the upper level. And for Type 5, the PT anomaly within the PBL is negative, while that above the PBL top is positive. In contrast, the PT anomaly of other types (2, 3, 6, and 7) tends to weaken the thermal inversion at the PBL top, which would favor the growth of the PBL to some extent.

How do these different PBL thermal structures develop under different synoptic conditions? As discussed in Sect. 3.2, the high CLD could reduce the solar radiation reaching the land surface, leading to suppression of the PBL during the daytime. The cloudiness may be one factor leading to the different PBL structures. In addition, the BLH of different synoptic patterns is also highly related to RH2 and southerly winds at the 925 hPa level (Fig. 8). Can the different synoptic winds and resultant warm/cold advectations modulate the PBL structure?

To determine the possible effects of warm/cold advection on PBL structure, the three-dimensional PT fields and wind vectors obtained from the idealized simulations were analyzed, in which the effects of cloudiness were isolated. As with the simulated BLH at 14:00 BJT shown in Fig. 10, significant differences could be noted among the simulated BLHs of different synoptic patterns. In Fig. 10, the simulated BLH at 14:00 BJT of each synoptic pattern is also



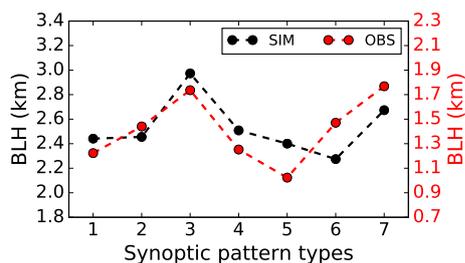
**Figure 8.** Correlations ( $R$ ) between the mean values of  $\text{PM}_{2.5}$  concentration and meteorological parameters for the different synoptic patterns, including (from left to right) 2 m temperature (T2), 2 m relative humidity (RH2), wind speed at the 925 hPa level (WS), south- and north-wind frequencies at the 925 hPa level (WD), total cloud cover at 14:00 BJT (CLD), and the BLH at 14:00 BJT. The grey bars represent the correlations between BLH and these meteorological parameters. Bars outlined in thick black lines indicate correlation coefficients ( $R$ ) that are statistically significant ( $p < 0.05$ ). Note that the  $R$  is calculated based on the seven pairs of mean values for different synoptic patterns.



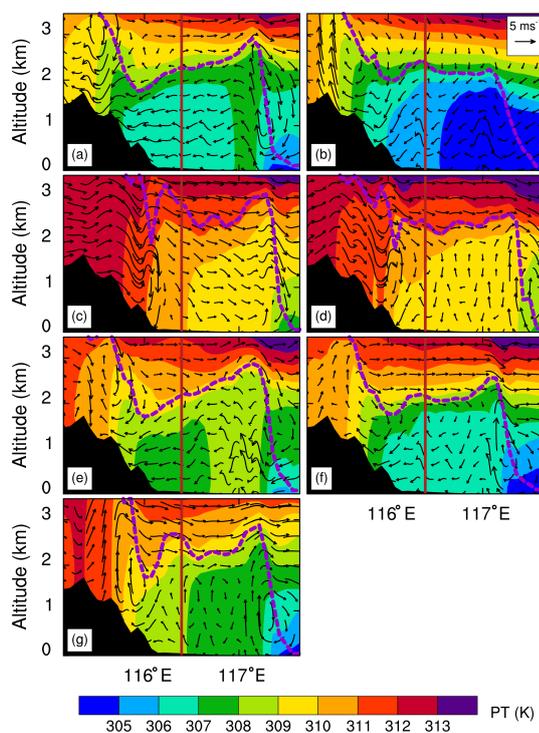
**Figure 9.** Vertical profiles of potential temperature (PT) at 14:00 BJT associated with the seven types of synoptic pattern derived from soundings (red lines). Solid lines indicate average values and shaded areas show the uncertainty range (the mean  $\pm 1$  standard deviation). Green solid lines represent the summer averaged BLH and green dashed lines represent the average BLH for each synoptic type. The PT anomaly (subtracted from the summer averaged PT profile) for each synoptic type was also given by the blue dashed lines in each panel.

compared with that derived from soundings. Both the simulations and observations show the same variation trend of BLH from Type 1 to Type 7, except for the relatively high BLH of Type 6 simulated by WRF. In other words, the variations of BLH among the seven synoptic patterns are well reproduced by the idealized simulations, although the simulated BLH is generally higher than the observed BLH, which may be caused by the simplification of cloud processes in

the idealized simulations. With the weakest pressure gradient over Beijing (Fig. 5f), the large-scale synoptic forcing of Type 6 to the Beijing area may be overwhelmed by the local atmospheric circulations (e.g., sea-breeze circulation and plain–mountain breeze circulation) in the idealized simulations with intense insolation, which may lead to the failure simulation of the PBL structure for Type 6. Despite this discrepancy, the simulated BLH variations of another six syn-



**Figure 10.** The simulated BLH at 14:00BJT (in black) and observed BLH at 14:00BJT (in red) over Beijing were shown as a function of synoptic types.



**Figure 11.** Vertical cross sections of the PT and wind vectors at 14:00BJT for the seven synoptic pattern types: (a) Type 1, (b) Type 2, (c) Type 3, (d) Type 4, (e) Type 5, (f) Type 6, and (g) Type 7. The vertical wind component is multiplied by a factor of 30 when plotting the wind vector fields. The black area on the left side of each panel shows the terrain. The red line shows the location of the sounding station in Beijing. The violet dashed line marks the BLH estimated using the *Ri* method. The location of the cross section is shown by the red line in Fig. 5a.

optic patterns show good agreement with the variations of observed BLH, which provides a basis for using the simulation results of these six types to unravel the possible mechanisms affecting PBL structure. For these six synoptic types, the simulated BLH of Types 3 and 7 is significantly higher than that of Types 1, 2, 4, and 5 (Fig. 10).

Figure 11 shows the cross sections of simulated PT from Beijing to the coastline of the Bohai Sea. For Types 1 and 2,

the southeasterly and easterly PBL winds could bring the cool marine air from the Bohai Sea to Beijing, leading to cooling of the PBL over Beijing in the afternoon (Figs. 11a–b, 6a, and S5). Such cold advections associated with these two synoptic types could be partially responsible for the observed cool anomalies within the PBL (Fig. 9a and b). And when the cooling of the PBL strengthens the thermal contrast between the PBL (cold) and its upper level (warm), the growth of the PBL could be suppressed, leading to exacerbation of the air pollution there.

In contrast, warm anomalies are found in the afternoon within the lower part of the troposphere for Types 3 and 4 (Fig. 9c and d), which may be caused by the westerly warm advections from the mountains (Figs. 11c–d and S6). Since the mountains act as elevated heat sources in the afternoon and warm the adjacent air (Fig. S6), in the presence of westerly winds above the mountains, the warmer air could be transported to the plain regions of Beijing (Figs. 11c–d and S6), imposing the warm anomalies on the PBL there. When the warm anomaly is stronger above the PBL than within the PBL, the development of the PBL could be suppressed.

For Type 5, the cold advection within the PBL and the warm advection aloft develop simultaneously over Beijing in the afternoon (Fig. 11e), strengthening the thermal inversion, capping the PBL top and suppressing the growth of the PBL. Such a combination of cold and warm advections could be partly responsible for the relatively low BLH associated with Type 5 (Fig. 9e). The simultaneous presence of cold advection and warm advection could also be found in the simulated cross section of Type 7, but it is less prominent. As illustrated in Figs. 5g and 6a, under the synoptic conditions of Type 7, the synoptic forcing (northerly winds) within the PBL do not support the advection of marine air from the Bohai Sea to Beijing; the simulated cold advection to Beijing is more likely to be induced by the local sea-breeze circulation.

In short, in addition to cloudiness, the results of idealized simulations demonstrate that the cold/warm advection induced by the synoptic forcing can also modulate the PBL structure in Beijing in the afternoon. With favorable thermal conditions in the summertime, the local atmospheric circulations (e.g., mountain–plain breeze, sea–land breeze) may be in place and superimposed on the large-scale synoptic advection, influencing the PBL structure in Beijing (Hu et al., 2014; Miao et al., 2015b).

## 4 Conclusions

In this study, seven different synoptic patterns during the summer in Beijing were identified using the T-PCA method and NCEP-FNL reanalysis data from 2011 to 2014. Their relationships with aerosol pollution and the PBL structure were comprehensively investigated using collocated long-

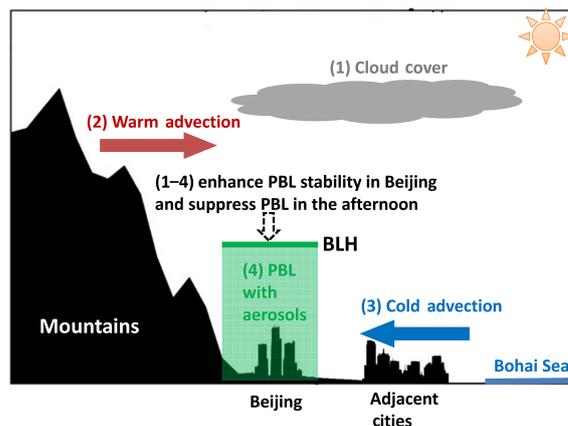
term PM<sub>2.5</sub> measurements and fine-resolution soundings in Beijing.

The climatological diurnal cycle of the BLH, revealed by thrice-daily soundings (at 08:00, 14:00, and 20:00 BJT), shows that the BLH reaches a maximum in the afternoon and is anti-correlated with the near-surface PM<sub>2.5</sub> concentration. In addition, daily aerosol pollution is significantly anti-correlated with the BLH at 14:00 BJT. The correlation analysis between the BLH and local meteorological parameters shows that in the afternoon, the BLH is negatively correlated with CLD and RH2. Thus, heavy aerosol pollution events frequently occur under cloudy conditions with high RH2 and low BLH.

By classifying the summertime 925 hPa GH fields over the Beijing–Tianjin–Hebei region (43–47° N, 112–125° E), three types of synoptic patterns favoring the occurrence of heavy aerosol pollution were identified, accounting for 67 % of the total. At the 925 hPa level, these three synoptic types were characterized by high-pressure systems located to the east or southeast of Beijing, leading to southerly PBL winds in Beijing. These southerly PBL winds favor the transport of pollutants from the surrounding southern industrial cities (e.g., Baoding, Shijiazhuang, and Cangzhou) to Beijing. Therefore, the horizontal transport of pollutants induced by the synoptic forcing may be one of the most important factors affecting the air quality of Beijing in summer. In the vertical dimension, these three synoptic types were characterized by a relatively low BLH at 14:00 BJT, which may be highly associated with the relatively high CLD. The presence of clouds would reduce the solar radiation reaching the land surface, and then suppress the development of the PBL during the daytime. The similar impacts on the surface solar radiation can be induced by aerosols within the PBL. Besides, the absorbing aerosols could heat the atmosphere, leading to a more stable and shallower PBL during the daytime.

In addition to the cloudiness and aerosols within the PBL, the large-scale atmospheric advection within the PBL may play an important role in modulating the development of the PBL, which was examined by idealized simulations. In the afternoon, cold advection from the south within the PBL, induced by large-scale synoptic forcing, could decrease the PBL temperature and strengthen the thermal inversion above the PBL, thus suppressing the development of the PBL. Moreover, when warm advection appears simultaneously at the top level of the PBL, the suppression of the PBL is intensified. The suppression of the PBL induced by the large-scale atmospheric advection under certain synoptic types, therefore, may also play a role in modulating the development of the PBL in Beijing. These factors/processes affecting the daytime PBL over Beijing in the afternoon are summarized by a schematic diagram in Fig. 12.

Although this study focuses on the daytime PBL structure, the nocturnal PBL also significantly affects the air quality at hourly to diurnal scales through the intermittent turbulence, which also cannot be ignored. The structure of the



**Figure 12.** Schematic diagram of factors/processes suppressing the PBL over Beijing in the afternoon, including (1) high cloud cover, (2) westerly warm advection above the PBL, (3) southerly cold advection within the PBL, and (4) aerosols within the PBL.

nocturnal PBL is primarily determined by Stull (1988) and Salmond and McKendry (2005). The nocturnal PBL may range from fully turbulent to intermittently turbulent or even non-turbulent at a variety of heights, temporal scales, and spatial locations, which was largely induced by complex interactions between the static stability of the atmosphere and those processes (i.e., wind shear from synoptic patterns, terrain-induced flows, low-level jets) that govern mechanical generation of turbulence. This makes it very difficult for the observation of large-scale atmospheric advection, investigation of the PBL–synoptic pattern interaction, transport pathways, and dispersion of pollutants in the nocturnal PBL, particularly in regions of complex terrain such as Beijing. To fully understand the impacts of the PBL on air quality in Beijing, more attention should be paid to the nocturnal PBL in the future.

Besides, although the important impacts of large-scale synoptic forcing on PBL structures and air quality in Beijing during the summer have been emphasized in this study, the important roles of local atmospheric circulations in the modulation of the PBL structure and air pollution cannot be overlooked. The presence of aerosols may also play a role in modifying PBL structures and processes, which warrants further study and will form the basis of our future work.

*Data availability.* The NCEP–FNL reanalysis and aerosol measurements are publicly available at <http://rda.ucar.edu/datasets/ds083.2/>, and [www.stateair.net](http://www.stateair.net), respectively. The meteorological observations, BLH data, and simulation results are available upon request via email: [jpguocams@gmail.com](mailto:jpguocams@gmail.com) (J. Guo).

**The Supplement related to this article is available online at doi:10.5194/acp-17-3097-2017-supplement.**

*Competing interests.* The authors declare that they have no conflict of interest.

*Acknowledgements.* This study is supported by the National Natural Science Foundation of China (91544217 and 41471301), the Ministry of Science and Technology of China (2014BAC16B01, 2013CB955804), and the Chinese Academy of Meteorological Sciences (2014R18). The authors would like to acknowledge the China Meteorological Administration for providing the long-term sounding and total cloud cover data, and the US diplomatic missions for providing the PM<sub>2.5</sub> concentration measurements. Last but not least, special thanks go to three anonymous reviewers for their valuable comments and suggestions that helped greatly to improve the quality of our manuscript.

Edited by: J. Huang

Reviewed by: three anonymous referees

## References

- Abdi, H. and Williams, L. J.: Principal component analysis, *Wiley Interdiscip. Rev. Comput. Stat.*, 2, 433–459, doi:10.1002/wics.101, 2010.
- Bernaards, C. A. and Jennrich, R. I.: Gradient projection algorithms and software for arbitrary rotation criteria in factor analysis, *Educ. Psychol. Meas.*, 65, 676–696, doi:10.1177/0013164404272507, 2005.
- Chan, C. K. and Yao, X.: Air pollution in mega cities in China, *Atmos. Environ.*, 42, 1–42, doi:10.1016/j.atmosenv.2007.09.003, 2008.
- Chen, F. and Dudhia, J.: Coupling an Advanced Land Surface–Hydrology Model with the Penn State–NCAR MM5 Modeling System, Part I: Model Implementation and Sensitivity, *Mon. Weather Rev.*, 129, 569–585, doi:10.1175/1520-0493(2001)129<0587:CAALSH>2.0.CO;2, 2001.
- Chen, Y., Zhao, C., Zhang, Q., Deng, Z., Huang, M., and Ma, X.: Aircraft study of mountain chimney effect of Beijing, China, *J. Geophys. Res.-Atmos.*, 114, 1–10, doi:10.1029/2008JD010610, 2009.
- Chen, Z. H., Cheng, S. Y., Li, J. B., Guo, X. R., Wang, W. H., and Chen, D. S.: Relationship between atmospheric pollution processes and synoptic pressure patterns in northern China, *Atmos. Environ.*, 42, 6078–6087, doi:10.1016/j.atmosenv.2008.03.043, 2008.
- Chung, A., Chang, D. P. Y., Kleeman, M. J., Perry, K. D., Cahill, T. A., Dutcher, D., McDougall, E. M., and Stroud, K.: Comparison of real-time instruments used to monitor airborne particulate matter, *J. Air Waste Manage. Assoc.*, 51, 109–120, doi:10.1080/10473289.2001.10464254, 2001.
- Craven, J. P., Jewell, R. E., and Brooks, H. E.: Comparison between observed convective cloud-base heights and lifting condensation level for two different lifted parcels, *Weather Forecast.*, 885–890, doi:10.1175/1520-0434(2002)017<0885:CBOCCB>2.0.CO;2, 2002.
- De Wekker, S. F. J.: Observational and numerical evidence of depressed convective boundary layer heights near a mountain base, *J. Appl. Meteorol. Climatol.*, 47, 1017–1026, doi:10.1175/2007JAMC1651.1, 2008.
- Ding, A. J., Huang, X., Nie, W., Sun, J. N., Kerminen, V.-M., Petäjä, T., Su, H., Cheng, Y. F., Yang, X.-Q., Wang, M. H., Chi, X. G., Wang, J. P., Virkkula, A., Guo, W. D., Yuan, J., Wang, S. Y., Zhang, R. J., Wu, Y. F., Song, Y., Zhu, T., Zilitinkevich, S., Kulmala, M., and Fu, C. B.: Black carbon enhances haze pollution in megacities in China, *Geophys. Res. Lett.*, 43, 1–7, doi:10.1002/2016GL067745, 2016.
- Garratt, J.: Review: the atmospheric boundary layer, *Earth-Science Rev.*, 37, 89–134, doi:10.1016/0012-8252(94)90026-4, 1994.
- Gao, Y., Zhang, M., Liu, Z., Wang, L., Wang, P., Xia, X., Tao, M., and Zhu, L.: Modeling the feedback between aerosol and meteorological variables in the atmospheric boundary layer during a severe fog–haze event over the North China Plain, *Atmos. Chem. Phys.*, 15, 4279–4295, doi:10.5194/acp-15-4279-2015, 2015.
- Guo, J. P., Zhang, X. Y., Wu, Y. R., Zhaxi, Y., Che, H. Z., La, B., Wang, W., and Li, X. W.: Spatio-temporal variation trends of satellite-based aerosol optical depth in China during 1980–2008, *Atmos. Environ.*, 45, 6802–6811, doi:10.1016/j.atmosenv.2011.03.068, 2011.
- Guo, J., Niu, T., Wang, F., Deng, M., and Wang, Y.: Integration of multi-source measurements to monitor sand-dust storms over North China: A case study, *Acta Meteorol. Sin.*, 27, 566–576, doi:10.1007/s13351-013-0409-z, 2013.
- Guo, J., Miao, Y., Zhang, Y., Liu, H., Li, Z., Zhang, W., He, J., Lou, M., Yan, Y., Bian, L., and Zhai, P.: The climatology of planetary boundary layer height in China derived from radiosonde and reanalysis data, *Atmos. Chem. Phys.*, 16, 13309–13319, doi:10.5194/acp-16-13309-2016, 2016.
- Guo, S., Hu, M., Zamora, M. L., Peng, J., Shang, D., Zheng, J., Du, Z., Wu, Z., Shao, M., Zeng, L., Molina, M. J., and Zhang, R.: Elucidating severe urban haze formation in China, *P. Natl. Acad. Sci. USA*, 111, 17373–17378, doi:10.1073/pnas.1419604111, 2014.
- Han, T., Liu, X., Zhang, Y., Qu, Y., Zeng, L., Hu, M., and Zhu, T.: Role of secondary aerosols in haze formation in summer in the Megacity Beijing, *J. Environ. Sci.*, 31, 51–60, doi:10.1016/j.jes.2014.08.026, 2015.
- He, K., Yang, F., Ma, Y., Zhang, Q., Yao, X., Chan, C. K., Cadle, S., Chan, T., and Mulawa, P.: The characteristics of PM<sub>2.5</sub> in Beijing, China, *Atmos. Environ.*, 35, 4959–4970, doi:10.1016/S1352-2310(01)00301-6, 2001.
- Hong, S.-Y., Noh, Y., and Dudhia, J.: A New Vertical Diffusion Package with an Explicit Treatment of Entrainment Processes, *Mon. Weather Rev.*, 134, 2318–2341, doi:10.1175/MWR3199.1, 2006.
- Hu, X., Ma, Z., Lin, W., Zhang, H., Hu, J., Wang, Y., Xu, X., Fuentes, J. D., and Xue, M.: Impact of the Loess Plateau on the atmospheric boundary layer structure and air quality in the North China Plain?: A case study, *Sci. Total Environ.*, 499, 228–237, doi:10.1016/j.scitotenv.2014.08.053, 2014.
- Huang, R.-J., Zhang, Y., Bozzetti, C., Ho, K.-F., Cao, J.-J., Han, Y., Daellenbach, K. R., Slowik, J. G., Platt, S. M., Canonaco, F., Zotter, P., Wolf, R., Pieber, S. M., Brun, E. A., Crippa, M., Ciarelli, G., Piazzalunga, A., Schwikowski, M., Abbaszade, G., Schnelle-Kreis, J., Zimmermann, R., An, Z., Szidat, S., Baltensperger, U., Haddad, I. El, and Prévôt, A. S. H.: High secondary aerosol contribution to particulate pollution during haze events in China, *Nature*, 514, 218–222, doi:10.1038/nature13774, 2014.

- Huth, R.: An intercomparison of computer-assisted circulation classification methods, *Int. J. Climatol.*, 16, 893–922, doi:10.1002/(SICI)1097-0088(199608)16:8<893::AID-JOC51>3.0.CO;2-Q, 1996.
- Huth, R.: A circulation classification scheme applicable in GCM studies, *Theor. Appl. Climatol.*, 67, 1–18, doi:10.1007/s007040070012, 2000.
- Huth, R., Beck, C., Philipp, A., Demuzere, M., Ustrnul, Z., Cahynová, M., Kyselý, J., and Tveito, O. E.: Classifications of atmospheric circulation patterns: Recent advances and applications, *Ann. Ny. Acad. Sci.*, 1146, 105–152, doi:10.1196/annals.1446.019, 2008.
- Iacono, M. J., Delamere, J. S., Mlawer, E. J., Shephard, M. W., Clough, S. A., and Collins, W. D.: Radiative forcing by long-lived greenhouse gases: Calculations with the AER radiative transfer models, *J. Geophys. Res.*, 113, D13103, doi:10.1029/2008JD009944, 2008.
- Kim, Y., Kim, S.-W., and Yoon, S.-C.: Observation of new particle formation and growth under cloudy conditions at Gosan Climate Observatory, Korea, *Meteorol. Atmos. Phys.*, 126, 81–90, doi:10.1007/s00703-014-0336-2, 2014.
- Kusaka, H., Kondo, H., Kikegawa, Y., and Kimura, F.: A Simple Single-Layer Urban Canopy Model For Atmospheric Models: Comparison With Multi-Layer And Slab Models, *Bound.-Lay. Meteorol.*, 101, 329–358, doi:10.1023/A:1019207923078, 2001.
- Li, J., Chen, H., Li, Z., Wang, P., Cribb, M., and Fan, X.: Low-level temperature inversions and their effect on aerosol condensation nuclei concentrations under different large-scale synoptic circulations, *Adv. Atmos. Sci.*, 32, 898–908, doi:10.1007/s00376-014-4150-z, 2015.
- Li, Z., Xia, X., Cribb, M., Mi, W., Holben, B., Wang, P., Chen, H., Tsay, S.-C., Eck, T. F., Zhao, F., Dutton, E. G., and Dickerson, R. E.: Aerosol optical properties and their radiative effects in northern China, *J. Geophys. Res.*, 112, 1–11, doi:10.1029/2006JD007382, 2007.
- Li, Z., Lau, W. K.-M., Ramanathan, V., Wu, G., Ding, Y., Manoj, M. G., Liu, J., Qian, Y., Li, J., Zhou, T., Fan, J., Rosenfeld, D., Ming, Y., Wang, Y., Huang, J., Wang, B., Xu, X., Lee, S.-S., Cribb, M., Zhang, F., Yang, X., Zhao, C., Takemura, T., Wang, K., Xia, X., Yin, Y., Zhang, H., Guo, J., Zhai, P. M., Sugimoto, N., Babu, S. S., and Brasseur, G. P.: Aerosol and monsoon climate interactions over Asia, *Rev. Geophys.*, 119–161, doi:10.1002/2015RG000500, 2016.
- Liu, S. and Liang, X. Z.: Observed diurnal cycle climatology of planetary boundary layer height, *J. Climate*, 23, 5790–5809, doi:10.1175/2010JCLI3552.1, 2010.
- Liu, S., Liu, Z., Li, J., Wang, Y., Ma, Y., Sheng, L., Liu, H., Liang, F., Xin, G., and Wang, J.: Numerical simulation for the coupling effect of local atmospheric circulations over the area of Beijing, Tianjin and Hebei Province, *Sci. China Ser. D*, 52, 382–392, doi:10.1007/s11430-009-0030-2, 2009.
- Liu, X. G., Li, J., Qu, Y., Han, T., Hou, L., Gu, J., Chen, C., Yang, Y., Liu, X., Yang, T., Zhang, Y., Tian, H., and Hu, M.: Formation and evolution mechanism of regional haze: a case study in the megacity Beijing, China, *Atmos. Chem. Phys.*, 13, 4501–4514, doi:10.5194/acp-13-4501-2013, 2013.
- Liu, X., Yan, L., Yang, P., Yin, Z. Y., and North, G. R.: Influence of Indian summer monsoon on aerosol loading in East Asia, *J. Appl. Meteorol. Climatol.*, 50, 523–533, doi:10.1175/2010JAMC2414.1, 2011.
- Liu, Z., Wang, Y., Hu, B., Ji, D., Zhang, J., Wu, F., Wan, X., and Wang, Y.: Source appointment of fine particle number and volume concentration during severe haze pollution in Beijing in January 2013, *Environ. Sci. Pollut. Res.*, 23, 6845–6860, doi:10.1007/s11356-015-5868-6, 2016.
- Miao, Y., Liu, S., Zheng, Y., Wang, S., and Chen, B.: Numerical study of the effects of topography and urbanization on the local atmospheric circulations over the Beijing-Tianjin-Hebei, China, *Adv. Meteorol.*, 2015, 1–16, doi:10.1155/2015/397070, 2015a.
- Miao, Y., Hu, X.-M., Liu, S., Qian, T., Xue, M., Zheng, Y., and Wang, S.: Seasonal variation of local atmospheric circulations and boundary layer structure in the Beijing-Tianjin-Hebei region and implications for air quality, *J. Adv. Model. Earth Syst.*, 7, 1–25, doi:10.1002/2015MS000522, 2015b.
- Miao, Y., Liu, S., Zheng, Y., and Wang, S.: Modeling the feedback between aerosol and boundary layer processes: a case study in Beijing, China, *Environ. Sci. Pollut. Res.*, 23, 3342–3357, doi:10.1007/s11356-015-5562-8, 2016.
- Miller, S. T. K.: Sea breeze: Structure, forecasting, and impacts, *Rev. Geophys.*, 41, 1011, doi:10.1029/2003RG000124, 2003.
- Niu, F., Li, Z., Li, C., Lee, K. H., and Wang, M.: Increase of wintertime fog in China: Potential impacts of weakening of the Eastern Asian monsoon circulation and increasing aerosol loading, *J. Geophys. Res.-Atmos.*, 115, 1–12, doi:10.1029/2009JD013484, 2010.
- Peng, J., Hu, M., Guo, S., Du, Z., Zheng, J., Shang, D., Levy Zamora, M., Zeng, L., Shao, M., Wu, Y.-S., Zheng, J., Wang, Y., Glen, C. R., Collins, D. R., Molina, M. J., and Zhang, R.: Markedly enhanced absorption and direct radiative forcing of black carbon under polluted urban environments, *P. Natl. Acad. Sci. USA*, 113, 4266–4271, doi:10.1073/pnas.1602310113, 2016.
- Philipp, A., Bartholy, J., Beck, C., Ericum, M., Esteban, P., Fettweis, X., Huth, R., James, P., Jourdain, S., Kreienkamp, F., Krennert, T., Lykoudis, S., Michalides, S. C., Pianko-Kluczynska, K., Post, P., Álvarez, D. R., Schiemann, R., Spekat, A., and Tymvios, F. S.: Cost733cat – A database of weather and circulation type classifications, *Phys. Chem. Earth*, 35, 360–373, doi:10.1016/j.pce.2009.12.010, 2010.
- Pu, B. and Dickinson, R. E.: Diurnal spatial variability of Great Plains summer precipitation related to the dynamics of the low-level jet, *J. Atmos. Sci.*, 71, 1807–1817, doi:10.1175/JAS-D-13-0243.1, 2014.
- Quan, J., Gao, Y., Zhang, Q., Tie, X., Cao, J., Han, S., Meng, J., Chen, P., and Zhao, D.: Evolution of planetary boundary layer under different weather conditions, and its impact on aerosol concentrations, *Particuology*, 11, 34–40, doi:10.1016/j.partic.2012.04.005, 2013.
- Quan, J., Tie, X., Zhang, Q., Liu, Q., Li, X., Gao, Y., and Zhao, D.: Characteristics of heavy aerosol pollution during the 2012–2013 winter in Beijing, China, *Atmos. Environ.*, 88, 83–89, doi:10.1016/j.atmosenv.2014.01.058, 2014.
- Richman, M. B.: Obliquely rotated principal components: An improved meteorological map typing technique?, *J. Appl. Meteorol.*, 20, 1145–1159, doi:10.1175/1520-0450(1981)020<1145:ORPCA1>2.0.CO;2, 1981.

- Salmond, J. A. and McKendry, I. G.: A review of turbulence in the very stable nocturnal boundary layer and its implications for air quality, *Prog. Phys. Geogr.*, 29, 171–188, doi:10.1191/0309133305pp442ra, 2005.
- San Martini, F. M., Hasenkopf, C. A., and Roberts, D. C.: Statistical analysis of PM<sub>2.5</sub> observations from diplomatic facilities in China, *Atmos. Environ.*, 110, 174–185, doi:10.1016/j.atmosenv.2015.03.060, 2015.
- Seidel, D. J., Zhang, Y., Beljaars, A., Golaz, J.-C., Jacobson, A. R., and Medeiros, B.: Climatology of the planetary boundary layer over the continental United States and Europe, *J. Geophys. Res.-Atmos.*, 117, 1–15, doi:10.1029/2012JD018143, 2012.
- Stefan, S., Necula, C., and Georgescu, F.: Analysis of long-range transport of particulate matters in connection with air circulation over Central and Eastern part of Europe, *Phys. Chem. Earth*, 35, 523–529, doi:10.1016/j.pce.2009.12.008, 2010.
- Stull, R. B. (Ed.): An introduction to boundary layer meteorology, Springer Netherlands, Dordrecht, 1988.
- Sun, Y., Song, T., Tang, G., and Wang, Y.: The vertical distribution of PM<sub>2.5</sub> and boundary-layer structure during summer haze in Beijing, *Atmos. Environ.*, 74, 413–421, doi:10.1016/j.atmosenv.2013.03.011, 2013.
- Tang, G., Zhang, J., Zhu, X., Song, T., Munkel, C., Hu, B., Schäfer, K., Liu, Z., Zhang, J., Wang, L., Xin, J., Suppan, P., and Wang, Y.: Mixing layer height and its implications for air pollution over Beijing, China, *Atmos. Chem. Phys.*, 16, 2459–2475, doi:10.5194/acp-16-2459-2016, 2016.
- Tie, X., Zhang, Q., He, H., Cao, J., Han, S., Gao, Y., Li, X., and Jia, X. C.: A budget analysis of the formation of haze in Beijing, *Atmos. Environ.*, 100, 25–36, doi:10.1016/j.atmosenv.2014.10.038, 2015.
- Vogelezang, D. H. P. and Holtslag, A. A. M.: Evaluation and model impacts of alternative boundary-layer height formulations, *Bound.-Lay. Meteorol.*, 81, 245–269, doi:10.1007/BF02430331, 1996.
- Wang, F., Chen, D. S., Cheng, S. Y., Li, J. B., Li, M. J., and Ren, Z. H.: Identification of regional atmospheric PM<sub>10</sub> transport pathways using HYSPLIT, MM5-CMAQ and synoptic pressure pattern analysis, *Environ. Modell. Softw.*, 25, 927–934, doi:10.1016/j.envsoft.2010.02.004, 2010.
- Wang, H., Xu, J., Zhang, M., Yang, Y., Shen, X., Wang, Y., Chen, D., and Guo, J.: A study of the meteorological causes of a prolonged and severe haze episode in January 2013 over central-eastern China, *Atmos. Environ.*, 98, 146–157, doi:10.1016/j.atmosenv.2014.08.053, 2014.
- Wang, L., Liu, Z., Sun, Y., Ji, D., and Wang, Y.: Long-range transport and regional sources of PM<sub>2.5</sub> in Beijing based on long-term observations from 2005 to 2010, *Atmos. Res.*, 157, 37–48, doi:10.1016/j.atmosres.2014.12.003, 2015a.
- Wang, L., Xin, J., Li, X., and Wang, Y.: The variability of biomass burning and its influence on regional aerosol properties during the wheat harvest season in North China, *Atmos. Res.*, 157, 153–163, doi:10.1016/j.atmosres.2015.01.009, 2015b.
- Wei, P., Cheng, S., Li, J., and Su, F.: Impact of boundary-layer anticyclonic weather system on regional air quality, *Atmos. Environ.*, 45, 2453–2463, doi:10.1016/j.atmosenv.2011.01.045, 2011.
- Wilde, N. P., Stull, R. B., and Eloranta, E. W.: The LCL zone and cumulus onset, *J. Clim. Appl. Meteorol.*, 24, 640–657, doi:10.1175/1520-0450(1985)024<0640:TLZACO>2.0.CO;2, 1985.
- Wu, G. X., Li, Z. Q., Fu, C. B., Zhang, X. Y., Zhang, R. Y., Zhang, R. H., Zhou, T. J., Li, J. P., Li, J. D., Zhou, D. G., Wu, L., Zhou, L. T., He, B., and Huang, R. H.: Advances in studying interactions between aerosols and monsoon in China, *Sci. China Earth Sci.*, 59, 1–16, doi:10.1007/s11430-015-5198-z, 2016.
- Ye, X., Song, Y., Cai, X., and Zhang, H.: Study on the synoptic flow patterns and boundary layer process of the severe haze events over the North China Plain in January 2013, *Atmos. Environ.*, 124, 129–145, doi:10.1016/j.atmosenv.2015.06.011, 2016.
- Yoo, J.-M., Lee, Y.-R., Kim, D., Jeong, M.-J., Stockwell, W. R., Kundu, P. K., Oh, S.-M., Shin, D.-B., and Lee, S.-J.: New indices for wet scavenging of air pollutants (O<sub>3</sub>, CO, NO<sub>2</sub>, SO<sub>2</sub>, and PM<sub>10</sub>) by summertime rain, *Atmos. Environ.*, 82, 226–237, doi:10.1016/j.atmosenv.2013.10.022, 2014.
- Zhang, Y. L. and Cao, F.: Fine particulate matter (PM<sub>2.5</sub>) in China at a city level, *Sci. Rep.*, 5, 14884, doi:10.1038/srep14884, 2015.
- Zhang, J. P., Zhu, T., Zhang, Q. H., Li, C. C., Shu, H. L., Ying, Y., Dai, Z. P., Wang, X., Liu, X. Y., Liang, A. M., Shen, H. X., and Yi, B. Q.: The impact of circulation patterns on regional transport pathways and air quality over Beijing and its surroundings, *Atmos. Chem. Phys.*, 12, 5031–5053, doi:10.5194/acp-12-5031-2012, 2012.
- Zhang, L., Liao, H., and Li, J.: Impacts of Asian summer monsoon on seasonal and interannual variations of aerosols over eastern China, *J. Geophys. Res.*, 115, 1–21, doi:10.1029/2009JD012299, 2010.
- Zhang, R., Jing, J., Tao, J., Hsu, S.-C., Wang, G., Cao, J., Lee, C. S. L., Zhu, L., Chen, Z., Zhao, Y., and Shen, Z.: Chemical characterization and source apportionment of PM<sub>2.5</sub> in Beijing: seasonal perspective, *Atmos. Chem. Phys.*, 13, 7053–7074, doi:10.5194/acp-13-7053-2013, 2013.
- Zhang, R., Wang, G., Guo, S., Zamora, M. L., Ying, Q., Lin, Y., Wang, W., Hu, M., and Wang, Y.: Formation of Urban Fine Particulate Matter, *Chem. Rev.*, 115, 3803–3855, doi:10.1021/acs.chemrev.5b00067, 2015.
- Zhang, W., Guo, J., Miao, Y., Liu, H., Zhang, Y., Li, Z., and Zhai, P.: Planetary boundary layer height from CALIOP compared to radiosonde over China, *Atmos. Chem. Phys.*, 16, 9951–9963, doi:10.5194/acp-16-9951-2016, 2016.
- Zhang, Z., Zhang, X., Gong, D., Quan, W., Zhao, X., Ma, Z., and Kim, S. J.: Evolution of surface O<sub>3</sub> and PM<sub>2.5</sub> concentrations and their relationships with meteorological conditions over the last decade in Beijing, *Atmos. Environ.*, 108, 67–75, doi:10.1016/j.atmosenv.2015.02.071, 2015.
- Zhao, Y., Zhang, Q., Du, Y., Jiang, M., and Zhang, J.: Objective analysis of circulation extremes during the 21 July 2012 torrential rain in Beijing, *Acta Meteorol. Sin.*, 27, 626–635, doi:10.1007/s13351-013-0507-y, 2013.
- Zhu, P. and Albrecht, B.: A theoretical and observational analysis on the formation of fair-weather cumuli, *J. Atmos. Sci.*, 59, 1983–2005, doi:10.1175/1520-0469(2002)059<1983:ATAOAO>2.0.CO;2, 2002.
- Zhu, L., Huang, X., Shi, H., Cai, X., and Song, Y.: Transport pathways and potential sources of PM<sub>10</sub> in Beijing, *Atmos. Environ.*, 45, 594–604, doi:10.1016/j.atmosenv.2010.10.040, 2011.